

Borings are not boring: Examples of macrobioerosion in marine palaeo-ecosystems

Las perforaciones no aburren: Ejemplos de macrobioerosión en paleoecosistemas marinos

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Abstract: Bioerosion is an ecological process identifiable in the fossil record by means of traces left on hard substrates by producers, since Late Proterozoic times. It implies both the destruction and construction of information, and its analysis is a valuable tool to better understand the biodiversity and the (palaeo)ecological complexity of a given area at any moment of the Earth's history, as bioeroding organisms often lack a fossilisable hard skeleton. Moreover, bioerosion traces inform on two (reaction and coaction) of the three relationships established between organisms and their environment within the ecosystem. But bioerosion is also a first order taphonomic agent. Bioerosion and taphonomy follow one another as (palaeo)ecological and fossilisation processes, but at the same time they overlap and condition each other. This review aims to be a state-of-the-art treatise on bioerosion as both an ecological process and a taphonomic agent. The review mainly focuses on the fields of palaeomalacology, understanding molluscs both as producers and victims, or substrates of bioerosion, and of the rocky-shore environments, both centred in the Neogene period. Bioerosive studies of a palaeoecological and palaeoenvironmental nature are numerous and continue to increase. Research on bioerosion-taphonomy interaction still has a long way to go. As such, this contribution will demonstrate its value for both palaeoenvironmental interpretation and for understanding the sometimes-complex processes of fossilisation

Resumen: La bioerosión es un proceso ecológico identificable en el registro fósil mediante las trazas dejadas en sustratos duros por los productores, desde finales del Proterozoico. Implica tanto la destrucción como la construcción de información, y su análisis es una herramienta valiosa para entender mejor la biodiversidad y la complejidad (paleo) ecológica de una zona dada en cualquier momento de la historia de la Tierra, ya que los organismos bioerosivos a menudo carecen de un esqueleto duro fosilizable. Además, las trazas de bioerosión informan sobre dos de las tres posibles relaciones establecidas entre organismos y su entorno dentro del ecosistema. Pero la bioerosión también es un agente tafonómico de primer orden. La bioerosión y la tafonomía se suceden como procesos (paleo)ecológicos y de fosilización, pero al mismo tiempo se superponen y condicionan mutuamente. Esta revisión tiene como objetivo ser un tratado sobre la bioerosión como proceso ecológico y agente tafonómico. La revisión se centra básicamente en los campos de la paleomalacología, comprendiendo a los moluscos tanto como productores como víctimas o sustratos de bioerosión, y de los medios de costa rocosa, ambos centrados en el periodo Neógeno. Los estudios bioerosivos de naturaleza paleoecológica y paleoambiental son numerosos y continúan aumentando. La investigación sobre la interacción bioerosión-tafonomía todavía tiene un largo camino por recorrer. Esta contribución demostrará su valor tanto para la interpretación paleoambiental como para comprender los a veces complejos procesos de fosilización.

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INTRODUCTION

About fifty years ago I published my first scientific paper (Martinell, 1973), basically dedicated to the taphonomy and palaeoecology of a marine Pliocene outcrop in the northeast of the Iberian Peninsula. I included a chapter on the bioerosive activity of naticid gastropods. At that time, the bibliographic information on bioerosion was very scarce (Boekschooten, 1967). Currently, the

situation has changed radically and bioerosion is the subject of thematic congresses, books, and articles published in major scientific journals. In any case, there is still a strong bias in general palaeontology books, in which bioerosion is treated in a limited way at most, while bioturbation — the most traditional area of ichnology — receives all the attention. In the present

review I propose a brief tour of macrobioerosion, both from the palaeoecological and taphonomic fields, where it plays a leading role. For reasons related to my scientific background, I have limited this journey to the Cenozoic and mainly (although not exclusively) in relation to the position of molluscs as bioeroders or bioeroded agents.

The study of bioerosion is a valuable tool to understand the biodiversity and the ecological or palaeoecological complexity of a given area at any moment of the Earth's history. As bioeroding organisms often lack a fossilisable hard skeleton, it is the trace they leave on the substrate that enables determination of their existence in the past, and furthermore, to interpret the biotic interactions established with the substrate.

Bioerosion is a widespread ecological process that has transformed and modified habitats in all biomes of the Earth from the Late Proterozoic to Recent, by means of physical and/or chemical mechanisms. Much bioerosion takes place through the process of boring (also referred to as drilling) by organisms. Carbonates are the most common substrate into which they bore. Nevertheless, some examples are known from other lithologies (Masuda, 1968; Martinell, 1981; Ledesma-Vázquez & Johnson, 1994; Johnson *et al.*, 2010; Mayoral *et al.*, 2020). Wood is considered a special type of hard substrate (Noda & Lee, 1989; Evans, 1999; Taylor & Wilson, 2003). Bioeroders belong to numerous taxa including microbes (bacteria, algae, fungi, protists), lichens, vascular plants, invertebrates and vertebrates (Warme, 1975; Bromley, 1978; Ekdale *et al.*, 1984; Mayoral, 1988a, 1998b; Bromley, 2004; Glaub *et al.*, 2007; Cockell & Herrera, 2008). The structures produced by epiliths (organisms adapted to life upon a hard substrate) are primarily raspings, scrapings, superficial etching scars, and breakages; borings and bioclustrations are mainly due to endoliths (organisms that dwell *within* a hard substrate, which can be inorganic or organic in origin) (Neumann, 1966; Ekdale *et al.*, 1984; Martinell, 1989; Bromley, 1996; Gibert *et al.*, 2004; Wilson, 2007; Bertling *et al.*, 2007; Santos *et al.*, 2012a; Baarli *et al.*, 2017; Wisshak *et al.*, 2019).

Organisms and their environment form a reciprocating system, the ecosystem. Within it, three and only three types of interactions are possible: (1) living organisms that affect other living organisms (coaction), (2) living organisms that affect the abiotic environment (reaction), and (3) the inorganic environment that affects living organisms (action) (Lawrence, 1971; De Renzi, 1978). Ecology and palaeoecology must be logically defined as the study of these three interacting pairs, and bioerosion particularly is framed within the analysis of both coaction and reaction. Bioerosion is also a first order taphonomic agent that is feasible in all main phases of the process (Fig.1).

Researchers use numerous non-exclusive terms to describe the many processes by which organisms alter, break down, and erode hard substrates. These

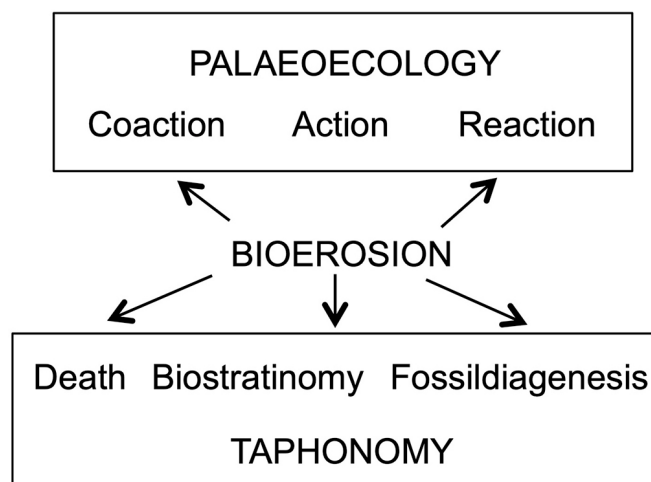


Figure 1. Graphical representation of formal relationships between bioerosion, taphonomy, and palaeoecology.

terms vary in function of the scope, scale, substrate type, and the discipline of study. For example, the erosive activities of fungi, lichens, and microbes on mineral substrates is called “biotic weathering” (Viles, 1995); the erosive effects of organisms on human-made substrates is termed “biodegradation” (Allsop *et al.*, 2004); the destruction or damage of a material due to a chemical reaction in front of an invasive organism is called “biocorrosion” (Beech & Gaylarde, 1999); “biogeomorphology” refers to the interactions and feedbacks between the landscape’s living and non-living components (Coombes, 2016); and the resulting cavity provoked by the growth-interfering behaviour of a symbiont that lives in the growing skeleton of the host is called “bioclustration” (Tapanila, 2004; Tapanila & Ekdale, 2007). The destruction and removal of consolidated mineral or lithic substrate by the direct action of organisms is referred to as “bioerosion” (Neumann, 1966). Indeed, the term itself has been reinterpreted, expanded, and/or redefined by multiple authors over the last half century (Ekdale *et al.*, 1984; Tapanila, 2008; Schönberg *et al.*, 2017). Nowadays, the term bioerosion is used by several authors to denote the breakdown of any hard substrate by biological activities, whether by microorganisms, invertebrates, vertebrates, or plants in marine, fresh water, or continental environments from the Late Proterozoic to Recent (Brasier *et al.*, 1994; Montalvo, 2002; Farlow & Holtz, 2002; Bromley, 2004; Hasiotis *et al.*, 2007; Mangano *et al.*, 2012; Davidson *et al.*, 2018; Wisshak *et al.*, 2019). More specifically, traces that are less than 1 mm produced by euendolithic microorganisms are referred as “microbioerosion” (Ekdale *et al.*, 1984; Glaub *et al.*, 2007; Wisshak, 2012) whereas “macrobioerosion” relates primarily to marine invertebrate bioeroders (McHuron, 1976; Ekdale *et al.*, 1984; Wilson, 2007; Wisshak *et al.*, 2019).

BIOEROSION VERSUS PALAEOECOLOGY

Reaction

Bioerosion in inorganic hard substrates. Bioeroders alter landscape features as a direct consequence of their activities and indirectly by accelerating erosional processes, causing substantial alterations to the morphology of landforms and habitats (Naylor *et al.*, 2012; Coombes, 2016; Domènech & Martinell, 2019). This activity is especially evident on the current and fossil carbonate rocky shores and on organic calcareous structures such as dead corals and shells. Macrobioerosion trace fossils are important indicators of intertidal and shallow subtidal marine palaeo-environments (Wilson & Palmer, 1990; Taylor & Wilson, 2003; Bromley, 2004; Santos *et al.*, 2011a; Schönberg & Wisshak, 2014). Rocks containing macroborings are not exclusive indicators of marine palaeo-ecosystems, but may also reflect freshwater habitats (Bolotov *et al.*, 2018, 2021).

The drilling activity of lithophagous bivalves in ancient cliffs had already been noted by Renaissance malacologists. According to Brocchi (1814), the Tertiary European pholadids had been cited by classical authors such as Aldrovandi, Targioni, Allioni, and Baldasieri, with studies based on the descriptive aspects of bivalves, but without delving into their boring activity.

The geological importance of the bioerosive activity of lithophagous molluscs was demonstrated as early as in Charles Lyell's *Principles of Geology* (Lyell, 1830), who started his seminal masterpiece with a reproduction of the abundant "*Lithodom*" borings found in the columns of the Temple of Serapis in Italy (currently known as Macellum of Pouzuoli), using them as a proof of relative sea level changes. Afterwards, Johnston (1850) classified the species responsible for those specific perforations as *Lithophaga lithophaga*.

Since the end of the 20th Century, bioerosive activity has been mainly used to characterize unconformable surfaces in the Palaeozoic (Pojeta & Palmer, 1976; James *et al.*, 1977; Kobluk *et al.*, 1977; Narbonne, 1984; Pemberton *et al.*, 1980; Johnson *et al.*, 1988; Desrochers & James, 1988; Webb, 1994; Rong & Johnson, 1996; Johnson *et al.*, 2010); the Mesozoic

(De La Beche, 1846; Kahrs, 1927; Hecker, 1965; Palmer & Fürsich, 1974; Bromley, 1975; Chudzikiewicz & Wieczorek, 1985; Ager, 1986; Gruszczynski, 1986; Breton *et al.*, 1992; Mikuláš, 1992; Wilson & Palmer, 1994; Cole & Palmer 1999; Farris *et al.*, 1999; Wilson *et al.*, 2005; Zatoń *et al.*, 2011; Krajewski *et al.*, 2017) and the Cenozoic (Lutzeier, 1922; Kiderlen, 1931; Lessertisseur, 1955; Aubert, 1972; Libbey & Johnson, 1997). It is clear nowadays that rocky-shore surfaces and their associated deposits are present throughout the entire geological register, commonly related to transgressive episodes (Johnson, 1988, 1992, 2006; Johnson & Baarli, 1999; Pemberton *et al.*, 2004; Santos *et al.*, 2008a; Santos & Mayoral, 2009). The ichnotaxonomic composition of rocky-shore assemblages changed throughout the Phanerozoic, although some taxa (*Trypanites*, *Gastrochaenolites*) are recurrent in the record. Figure 2 illustrates this evolution with three archetypical assemblages based on bibliographic data (see details in Gibert *et al.*, 2012). Neogene and Quaternary examples on fossil rocky-shore deposits are by far the most abundant and best documented ones (Pierçon, 1932; Denizot, 1952; Radwanski, 1970; Ballesio, 1972; Gutowski & Machalski, 1984; Rangheard, *et al.*, 1985; Watkins, 1990; Bromley & Asgaard, 1993a, 1993b; Uchman *et al.*, 2002; Titschack *et al.*, 2005; Charollais *et al.*, 2006; Johnson *et al.*, 2011; Demircan, 2012; Santos *et al.*, 2012b; Pineda-Salgado *et al.*, 2015; Brlek *et al.*, 2016; Abbel-Fattah & Assal, 2016; Santos & Mayoral, 2016; Steinthorsdottir & Håkansson, 2017; Hoffmann & Friedrich, 2017; El-Hedeny & El-Sabbagh, 2018; Martinell *et al.*, 2021b). In particular, there is a wealth of studies dealing with boring ichnoassemblages in Neogene rocky shores of the Iberian Peninsula (Martinell & Domènech, 1986, 1995; Silva *et al.*, 1995; Mayoral & Muñiz, 1996; Perry, 1996; Aguirre & Jiménez, 1997; Gibert *et al.*, 1998; Domènech *et al.*, 1999; Silva *et al.*, 1999; Domènech *et al.*, 2001; Cachão *et al.*, 2009; Santos *et al.*, 2008b, 2010, 2011b; Gibert *et al.*, 2012; Aguirre *et al.*, 2014; Rodríguez-Tovar *et al.*, 2015; Aguirre *et al.*, 2017; Giannetti, *et al.*, 2020).

The complex structures excavated by bioeroders can provide habitat for diverse and abundant

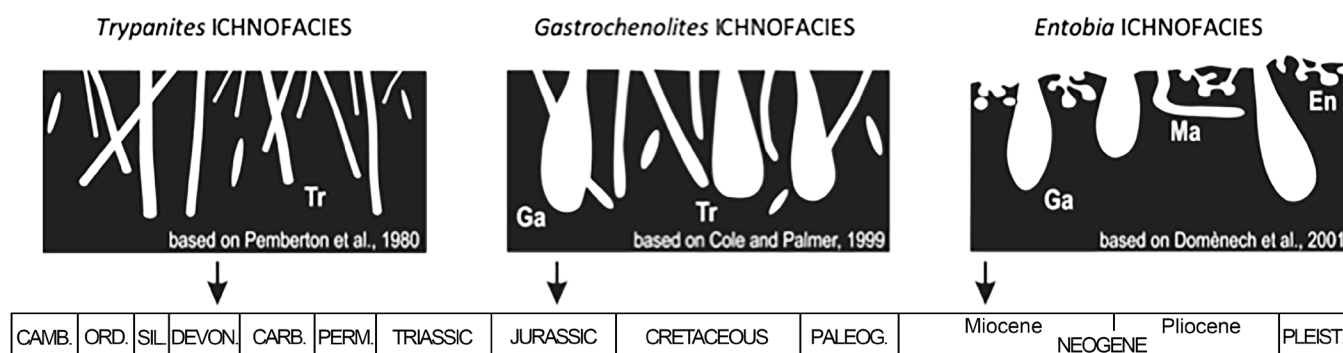


Figure 2. Evolution of the composition of rocky-shores trace fossil assemblages through the Phanerozoic, represented by three successive archetypical assemblages (see details in Gibert *et al.*, 2012).

species assemblages: recent communities inhabiting boreholes grated by bivalves in the rocky intertidal (mostly assigned to the ichnogenus *Gastrochaenolites* in the fossil record) contain 2.7 times more species than assemblages on bare rocks (Pinn *et al.*, 2008). However, intensive bioerosion can also impoverish biodiversity when such activity diminishes structural complexity or eliminates a habitat entirely.

Bioerosion of lithic substrates dominated by *Gastrochaenolites* indicates shallow marine environments. While lithophagous clams may also bore in deeper waters, a heavy infestation in a lithic substrate is largely a shallow water phenomenon (Kleemann, 1973).

As mentioned above, rocky shores in the fossil record appear associated with regional unconformities. The best palaeoenvironmental scenario for their formation and preservation occurs during major transgressive events capable of flooding rocky palaeoreliefs that are devoid of sedimentary cover. Marine evidence of submersion (bioerosion, body fossils, tidal or wave erosion) commonly overprints other features formed during the period of subaerial exposure (e.g., karstification in carbonate rocks) preceding the transgression (Bromley & Asgaard, 1993a, 1993b; Domènech *et al.*, 2001). Although it is possible to identify sedimentary deposits linked to rocky shores, these environments are essentially erosional, and bioerosion is often the best aid for their identification in the geological record. Bioeroded rocky-shore surfaces can be extensive and, given adequate outcrop conditions, can be recognized and correlated across different localities in a basin, thus contributing to its palaeogeographic reconstruction (Martinell *et al.*, 2021b). Such large surfaces are often time-averaged due to the overprinting of successive communities as the transgression proceeded. Bioerosion is also a very precise indicator of maximum sea level in those areas, since borings are observed to disappear above a certain altitude, a feature that can be related with other morphological characteristics that are indicative of sea-level limit, such as erosional notches. Identifying these palaeo sea-level positions may also be important to interpret eustatic changes, or relative changes due to tectonic activity (Gibert *et al.*, 1998, 2012; Domènech *et al.*, 2001; Cachão *et al.*, 2009; Giannetti *et al.*, 2020). Moreover, the density of coeval bioerosion combined with the estimation of the percentage of affected area allows one to precisely characterise the biological activity and to better calibrate palaeoenvironmental conditions (Domènech *et al.*, 2001; Martinell *et al.*, 2021b).

Regarding palaeoecological studies, rocky-shore assemblages offer the important advantage that organisms cementing to and boring into hard substrates generally retain their original positions on it during fossilisation, as represented by body fossils or by the traces themselves. In a transgressive situation, surfaces are exposed to deeper waters and new traces

overprint former ones. Later structures may cut across and destroy earlier ones, and endoliths can penetrate and modify previous structures (Bromley & Asgaard, 1993a; Domènech *et al.*, 2001; Gibert *et al.*, 2012; Silva *et al.*, 2019; Giannetti *et al.*, 2020). Thus, bioerosion itself leads to a masking or an elimination that does not always reflect the sequence of events (Bromley & Asgaard, 1993b; Domènech *et al.*, 2001). However, the morphologies of the borings that remain partial—those of the intact ones and those of the overlapped ones—have in some cases allowed estimation of the overall erosion on a linear scale. Thus, Domènech *et al.* (2001) proposed a denudation of several cm in the mid-marine Miocene wave-cutting platforms of Tarragona (Spain), based on the numerous and truncated *Gastrochaenolites* and *Phryxichnus* (both attributed to the action of lithophagous bivalves) present in them. The resulting surface exhibits *Entobia* (produced by endolithic but surficial boring sponges) and *Maeandropolydora* (due to the action of epilithic worms), which reveal a final stage of bioerosion (Fig. 3).

A transgressive situation often prevents identifying the ecological relationships that were established in the geological past for a given area. In the case of cliffs, erosion removes the skeletons of encrusters as well as many traces, so that only the most recent and/or deeper perforations can be observed. However, some examples exist where these relationships have been preserved. Thus, in the Jura marine Miocene, the ichnological record on Cretaceous limestone cliffs within the transgressive sequence has been studied. There are frequent dwelling buckets (*Circolites kotouensis* Mikuláš) of regular echinids, which contain numerous specimens of *Calulostrepsis* drilling that were produced by polychaetes (Bromley & D'Alessandro, 1983; Domènech *et al.*, 2008). Johnson *et al.* (2011) mentioned for the first time the presence of the ichnogenus *Caulostrepsis* in the concave surfaces of *Circolites* in the fossil record. These authors found rare specimens inside *C. kotouensis* in the Upper Miocene rocky shores on the north coast of Menorca (Balearic Islands). On the contrary, this association is well developed at the Les Bez outcrop. U-shaped annelid galleries are distributed in different patterns inside echinoid bowls, in some cases covering the entire surface, and in others being limited to their external rim. Fifty-five percent of the *C. kotouensis* at the site host *Caulostrepsis*, although the distribution is not regular, increasing towards the upper part of the cliff. This situation suggests an ecological relationship between the involved organisms—a relationship that has been preserved despite the erosional environment (Martinell *et al.*, 2021b). In the fossil record, *Caulostrepsis* has often been attributed to the spionid *Polydora* or the cirratulid *Dodecaceria*. At present, there is no documented relationship between *Polydora* and echinoid cavities, while different authors have highlighted this association with *Dodecaceria*

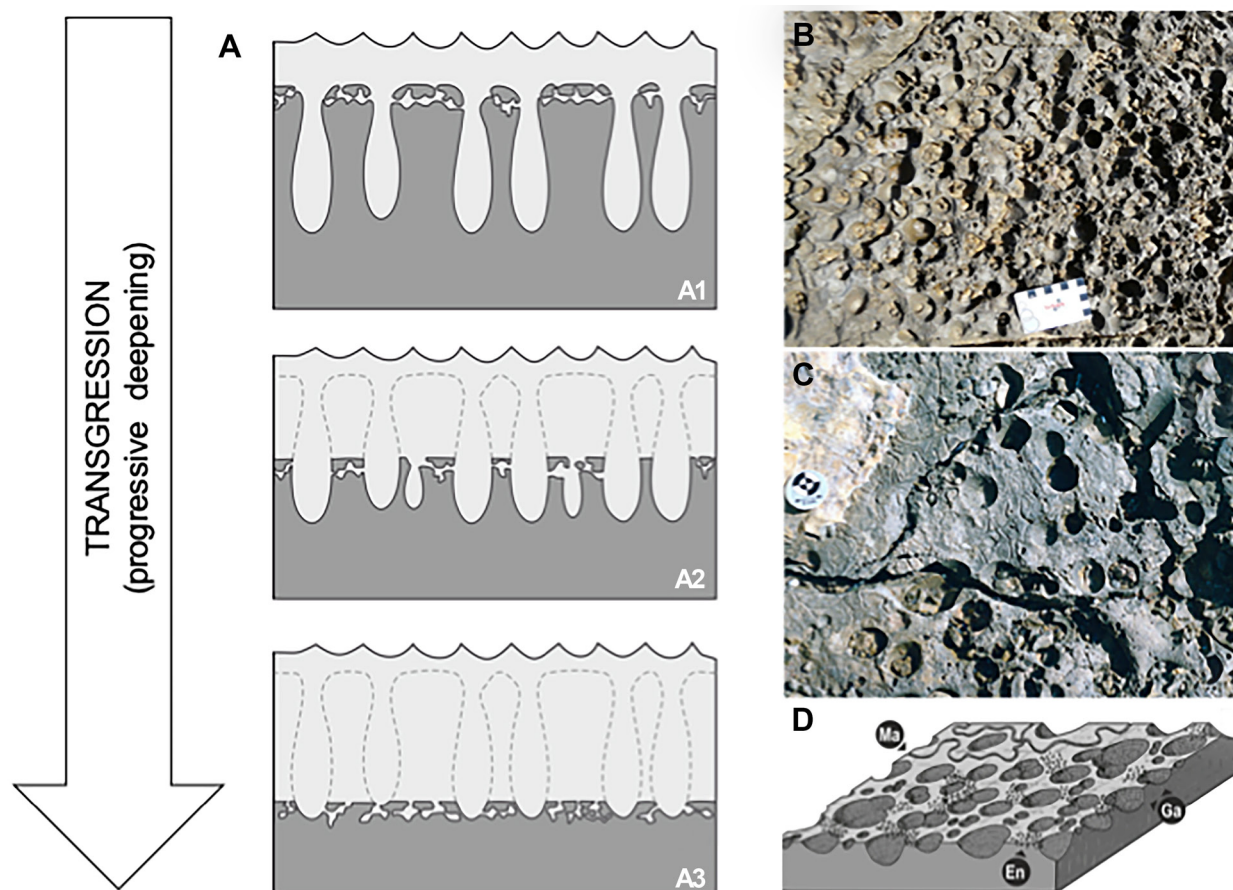


Figure 3. Bioerosive succession within the *Gastrochaenolites* Ichnocoenosis during a transgressive episode on a marine rocky-shore, Middle Miocene, Tarragona. **A**, From 1 to 3, evolution from the bivalve-dominated community (represented by *Gastrochaenolites*) to a sponge-dominated community (represented by *Entobia*); **B–C**, general overviews of the bored and abraded platform; **D**, Diagram of the final abrasion platform. **Ma**, *Maeandropolydora*; **Ga**, *Gastrochaenolites*; **En**, *Entobia*.

(Kempf, 1962; Voigt, 1970, 1971; Gibson, 2017; Gravina *et al.*, 2019). *Dodecaceria* infests encrusting calcareous algae, whereas regular echinoids feed on them. Therefore, in this case the association polychaete/echinoid could be explained by their mutual relationship to such algae. In several cases at Les Bez outcrop, the position of *Caulostrepsis* is limited to the distal zone of *Circolites*. That can be interpreted to be a result of the cohabitation between the producers, the echinoid offering passive protection to the worm against predators and/or the high energy environment (Davidson & Grupe, 2014).

Bioerosion in biological entities (organic remains). Skeletal remains (shells, bones, wood) deposited on the seabed act as islands for organisms that need a hard substrate to live. Although bioeroders may partially or totally destroy these remains, through their traces they also provide valuable information needed to reconstruct the biodiversity in ancient ecosystems. Known from the Late Proterozoic, species belonging to many groups of marine invertebrates and vertebrates have adopted a living habitat of boring, with up to 65 catalogued ichnogenera (Bromley, 2004; Wilson, 2013).

Sessile epilithic organisms attach themselves to the substrate in many ways. Permanently cementing species may secrete a carbonate skeleton in intimate contact with the substrate. When the organism requires a permanent location, an organic contact anchoring shallow superficial structures into a carbonate substrate provides the most durable bond. Such groups include verrucid barnacles (Darwin, 1854; Radwanski, 1977; Bromley & Martinell, 1991; Martinell & Domènech, 2022), eunicid polychaetes (Zibrowius, 1987; Martin & Britayev, 1998; Martinell & Domènech, 2009; Naimi *et al.*, 2021a), annelids (Santos *et al.*, 2003a), brachiopods (Bromley & Surlyk, 1973; Martinell, 1982a; Taddei Ruggiero & Annunziata, 2002; Jagt *et al.*, 2007; Bromley, 2008; Mergl, 2021), bryozoans (Mayoral, 1987a, 1988b; Mayoral & Reguant, 1995; Taylor *et al.*, 1999; Bromley & Heinberg, 2006), anomid bivalves (Bromley & Martinell, 1991; Taddei Ruggiero & Annunziata, 2002; Blissett & Pickeril, 2007; Neumann *et al.*, 2015) and vermetids (Mayoral, 1987b; Uchman *et al.*, 2017, 2018; Verde *et al.*, 2022).

Some groups of epilithic and endolithic organisms identifiable through their perforations have restricted environmental tolerances in terms of the bathymetric and sedimentary conditions. Invertebrate herbivores

that depend on epilithic and endolithic algae leave characteristic traces in the biological entities that indicate the photic zone. Therefore, the abundant presence of *Gnaticnus* (browsing and foraging traces attributed to dental erosion by regular echinoids) and/or *Radulichnus* (shallow fossil traces that have been attributed to the grazing activity of the radula of mollusks: chitons and gastropods) on hard surfaces such as the remains of shells or rocks indicate the epipelagic zone (Bromley, 1975; Voigt, 1977; Martinell & Domènech, 1981; Martinell, 1982b; Forrest *et al.*, 2001; Jagt, 2003; Gibert *et al.*, 2007; Cione *et al.*, 2010; Breton *et al.*, 2017; Lopes & Pereira, 2019).

Entobia (boring sponges), the most important ichnotaxon in advanced taphonomic stages of shell degradation, is abundant in both shallow and deep-water settings, as it seems to be particularly dependent on a sustained lack of sedimentation (De Groot, 1977; Wisshak, 2008; Garilli *et al.*, 2022). *Entobia* can involve the entire shell, resulting in stenomorphic borings (*sensu* Bromley & D'Alessandro, 1984, 1989). In this case, the destruction of the shell can be total (Belaústegui *et al.*, 2018). On certain surfaces, sponge borings are conspicuously absent, and worm borings (*Caulostrepsis*, *Trypanites*, etc.) predominate (Domènech *et al.*, 2008). These cases would represent situations in which sedimentation is problematic for some animals, being enough to prevent the colonization by sponges, but insufficient to prevent the proliferation of certain types of worms (Evans, 1969; Ekdale *et al.*, 1984; Hutchings, 2008).

Coaction

Traces of bioerosion present on skeletal remains could have been carried out during the life of the organism, indicating the cause of its death, or been produced *post-mortem*. The former ones correspond to predation or parasitism and commensalism (symbiosis), the latter are related to the activities of epizoa (endo- and epibionts) that use skeletal remains (shells, bones, wood) as a substrate. In this case, it is worth referring to it as taphonomic feedback (Kidwell & Jablonski, 1983). Proper recognition of shell damage is seminal since direct observations of predation are infeasible, but preserved shells with injuries are common in the fossil record.

Predation. The palaeontological marine record contains numerous examples of interspecific relationships (e.g., Tevesz & McCall, 1983; Boucot, 1990; Briggs & Crowther, 1990, 2003; Boucot & Poinar, 2010). Many of them correspond to exploitation pairs (predation and parasitism) among invertebrates, in which one member benefits at the expense of the other. The importance of predation as a selective influence on the evolution, diversification, mineral biomineralization and ecology of metazoans has been the focus of intensive research. Vermeij (1987), Kowalewski *et al.* (1998),

Conway-Morris and Bengtson (1994), Kowaleski and Kelly (2002), Leighton (2001, 2011), Kelley *et al.* (2003), Zuschin *et al.* (2003), and Martinell *et al.* (2010, 2012) are some examples of different approaches to the study of marine invertebrate predation in the fossil record.

The earliest evidence of predation in the fossil record comes from drill holes in *Cloudina* specimens from Ediacara strata of China (Bengtson & Zhao, 1992; Hua *et al.*, 2003), the oldest known skeletal fossil. Boring predators became geographically widespread during the Cambrian, having already attacked a variety of organisms (Kobluk *et al.*, 1978; Conway-Morris & Bengtson, 1994; Skovsted *et al.*, 2007). During the Late Cretaceous and Early Cenozoic, specialized marine durophagous taxa increased in diversity within the marine benthic ecosystems (Vermeij 1977; Bambach, 2002; Walker & Brett, 2002; Oji *et al.*, 2003; Harper, 2006; Zatoń & Salamon, 2008).

The Phanerozoic record of marine predation consists of both direct and indirect evidence. Indirect evidence is mostly based on predatory borings, shell repairs, skeletal regenerations, coprolites, and fossilized regurgitates (Conway Morris & Robinson, 1988; Baumiller & Gahn, 2004; Huntley & Kowalewski, 2007; Vallon, 2012; Zatoń & Rakocinski, 2014). Direct evidence in vertebrates corresponds to gut contents (Prikryl & Novosad, 2009; Zatoń *et al.*, 2017; Smit *et al.*, 2018; Vinn, 2018; Carnevale *et al.*, 2019; Hoffmann & Stevens, 2020).

In recent years, the concerted efforts of palaeontologists and ecologists have considerably advanced our knowledge of drilling predation and have demonstrated that drilling is a feeding mechanism employed by many groups of predators and that many shell-bearing groups of prey/host are drilled (Kabat, 1990; Kowalewski, 2002). In particular, the groups that have been most extensively studied as predators or hosts include molluscs (Vermeij, 1987; Dietl & Alexander, 2000; Kelley & Hansen, 2007; Martinell *et al.*, 2010; Naimi *et al.*, 2021b), brachiopods (Baumiller *et al.*, 1999; Taddei Ruggiero, 1999; Leighton, 2001; Harper, 2005; Kowalewski, 2002), and echinoderms (McClintock & Marion, 1993; Baumiller, 1996; Nebelsick & Kowalewski, 1999; Złotnik & Ceranka, 2005; Belaústegui *et al.*, 2017a).

Predators that drill holes in shells of marine invertebrates can provide quantitative data on various aspects of the predator-prey interactions, having produced one of the longest and richest fossil records in this scope (Palmer, 1982; Kitchell, 1986; Kelley & Hansen, 1993; Kowalewski *et al.*, 1998; Ceranka & Złotnik, 2003; Chattopadhyay & Baumiller, 2010).

Round drill-holes in shells made by predatory gastropods (dominantly, naticids and muricids) are among the few cases of direct evidence for behavioural interspecific interactions found in the fossil record. Such borings first appear in the Mesozoic and are very common in the Cenozoic fossil marine assemblages.

Morphologically, they can be classified into cylindrical (*Oichnus simplex* Bromley) and paraboloid borings (*Oichnus paraboloides* Bromley). The morphological consistency of fossil borings with those produced by Recent gastropods permits their approximate assignment to specific groups of predatory drilling gastropods (*O. simplex* to the muricids, *O. paraboloides* to the naticids) (Bromley, 1981; Hoffman & Martinell, 1984; Bromley, 2004). Although many other invertebrate groups produce similar borings, the reliability of the assignment can be further increased by the existence of associated body fossils of the suspected predators to the preys. However, it is prudent to consider other well-known producers and traces: Octopodid cephalopods that generate *O. ovalis* Bromley; eulimid gastropods that parasitise echinoderms and their traces, *O. halo* Nemann & Wisshak; or *Oichnus*-like holes of cassid gastropods in echinoids (Bromley, 1981; Neumann & Wisshak, 2009; Petsios *et al.*, 2021).

Traces of gastropod predation have been studied extensively at both Mesozoic and Cenozoic sites and for a variety of marine invertebrate preys, including annelids (Villegas-Martín *et al.*, 2016), ostracods (Reyment *et al.*, 1987; Reyment & Elewa, 2003; Baldanza *et al.*, 2020), barnacles (Gordillo, 2013; Klompmaker, *et al.*, 2015), decapod crustaceans (Pasini & Garassino, 2012; Klompmaker, *et al.*, 2013), brachiopods (Leighton, 2003; Harper, 2005; Baumiller *et al.*, 2006), bivalves and gastropods (Vermeij, 1982, 1994; Kabat, 1990; Dietl & Alexander, 2000; Kingsley-Smith *et al.*, 2003; Martinell *et al.*, 2010; Klompmaker, *et al.*, 2016), scaphopods (Yochelson *et al.*, 1983; Klompmaker, 2011), chitons (Rojas *et al.*, 2014), and echinoids (Neumann & Wisshak, 2009; Petsios *et al.*, 2021).

Naticids are the best studied group of molluscs regarding their predatory activity. Individually, they especially attack infaunal bivalves, but also other epifaunal molluscs—mainly gastropods—by means of drilling into their shells. The result is an O-shaped bore named *Oichnus paraboloides* (Bromley, 1981). The ability of naticids to move on or within the substrate allows them to easily locate their prey. Studies in aquariums allowed us to observe how they wrap the prey inside the foot in a mucus coat and then bury, drill and consume it completely in the sediment (unpublished personal observations). The naticids boring rate depends especially on the thickness and the ornamentation of the prey's shell (Reyment, 1967; Ansell, 1983; Kabat, 1990; Kabat, & Kohn, 1986; Kingsley-Smith *et al.*, 2003). Ziegelmeier (1954) and Kitchell *et al.* (1981) obtained similar drilling rates (0.6 mm/day) in Recent specimens of *Euspira nitida* Donovan and *Neverita duplicata* Say with an almost constant value regardless of prey species, predator size, or elapsed time.

Naticid gastropods are capable of an efficient selection of their prey. They preferably attack some of the potential prey species, select the size range of a particular prey depending on their own size, and selectively drill at

suitable areas of the prey shell. They are also expected to exclusively attack live individuals, the frequency of failed attacks (based on incomplete boreholes) being very low (Martinell, 1973; Martinell & Marquina, 1978; Kitchell *et al.*, 1981; Hoffman & Martinell, 1984; Kitchell *et al.*, 1986; Batllori & Martinell, 1992; Kingsley-Smith *et al.*, 2003; Kelley & Hansen, 2006). Prey can also detect the presence of naticids and try to escape predation. Observations on the behaviour of the bivalve *Callista chione* Linnaeus in terms of specimens of *Naticarius hebraeus* Martyn that were carried out in aquariums, showed that *C. chione* dug out and, with the help of its foot, moved around the aquarium when the naticids were introduced. Finally, being in a closed space, predators captured and subsequently consumed them (unpublished personal observations) (Fig. 4A–4C).

Naticids also feed on other naticids, but drillholes made by different species are almost identical. That makes it difficult to distinguish between confamilial, congeneric, or conspecific naticid predation based on the holes, and it is then better to consider it as confamilial predation rather than cannibalism among naticids (Fig. 4D). The confamilial predation among naticids has received notable attention from palaeoecologists and evolutionary palaeobiologists (e.g., Kelley, 1991; Dietl & Alexander, 2000; Kelley & Hansen, 2007; Fortunato, 2007; Martinell *et al.*, 2010; Mondal *et al.*, 2017). This is because confamilial predation among drilling gastropods provides quantitative and qualitative data in the context of important evolutionary and ecological concepts, including optimal foraging, escalation, size refuge, coevolution of dangerous prey, and behavioural ecology. The naticids' drilling patterns in confamilial depredation are similar to those observed in other prey. Although common, confamilial naticid predation intensity is not very prominent in the geological record (Kitchell *et al.*, 1981; Vermeij, 1982, 1994; Hoffman & Martinell, 1984; Kabat, 1990; Kelley, 1991; Kowalewski, 1993; Dietl & Alexander, 2000; Das *et al.*, 2014; Brezina *et al.*, 2016).

The Cretaceous radiation of naticids is part of the Mesozoic marine revolution, involving the increase in diversity of many modern marine predators due to the increase in shelled food supply (Palmer, 1982; Vermeij, 1977, 1982, 1987, 1994; Kabat, 1990; Kelley & Hansen, 1993; Walker & Brett, 2002; Bambach, 2002; Kase & Ishikawa, 2003).

The predatory activity of muricids is well known both in the Recent and in the fossil record. They attack epifaunal species of other gastropods, bivalves (especially ostreids and mytilids) and crustaceans (barnacles). Its predatory behaviour clearly differs from that of naticids. Their perforations are cylindrical and perpendicular to the shell thickness and the result is a conical-shaped borehole named *Oichnus simplex* (Bromley, 1981). Muricids can attack individually or in groups (causing several holes in the prey's shell), they neither select the location of the drilling that much, nor do they select the size of the prey, and they are rarely selective of

the prey itself (Bromley, 1981; Palmer, 1982; Dunkin & Hughes, 1984; Guerrero Alba & Reymont, 1988; Kowalewski, 2004; Yanes & Tyler, 2009; Martinelli *et al.*, 2013; Gordillo, 2013; Klompmaker *et al.*, 2015; Xu *et al.*, 2017).

The Muricidae evolutionary radiation preceded that of the balanomorphs during the Cretaceous. The evolutionary tendency of barnacles to reduce the number of their exoskeletal plates (from 8 to 4–6) since their appearance in the fossil record is probably a response to the predatory pressure of muricids that drilled into the imbrication zone of the plates (Palmer, 1982; Gale & Sørensen, 2015; Klompmaker *et al.*, 2015).

Echinoids are important prey for several vertebrate and invertebrate predatory groups in modern marine ecosystems. Carnivorous cassid gastropods are known to produce drill holes (*Oichnus*-like) in the process of preying upon echinoids, although other invertebrate groups have also been proposed as originators (Nebelsick & Kowalewski, 1999; Kowalewski & Nebelsick, 2003; Donovan & Pickerill, 2004; Złotnik & Ceranka, 2005; Meadows *et al.*, 2015). On the other hand, the parasitism of eulimid gastropods also produces drill holes, although significantly smaller, named *Oichnus halo* (Neumann & Wisshak, 2009).

The study of the distribution of the holes made by large and small *Cassis* (Gastropoda, Cassidae) on fossil echinoids from the Jurassic to the Recent verifies that, as in the case of naticids, the drilling activity was and is selective both for the size and place. Large cassids bore into and consume their prey almost always individually, whereas small cassids sometimes prey in groups upon urchins. Large cassids also display a higher site-selectivity. They more frequently bore into the petals and the apical disc. Larger cassids prey more habitually on larger echinoids (Farrar *et al.*, 2020). However, many trace morphologies observed on fossil echinoids cannot be associated with a possible predator. A better understanding of the morphology and distribution of these traces will allow us to carry out evaluations of the ecological and evolutionary importance of biotic interactions in benthic marine ecosystems (Złotnik & Ceranka, 2005; Tyler *et al.*, 2018; Farrar *et al.*, 2020; Petsios *et al.*, 2021).

Several modern octopod species belonging to the widespread genera *Octopus* and *Eledone* are known to drill into at least some of their shelled molluscan and crustacean prey. Drill holes made by octopods vary considerably in morphology, in part related to the microstructure and mineralogy of the affected shell (Wodinsky, 1973; Boyle & Knobock, 1981; Nixon & Maconnachie, 1988; McQuaid, 1994; Smith, 2003; López-Uriarte *et al.*, 2008; Fee *et al.*, 2023). All the first authors describe *Octopus* borings as oval. Wodinsky (1969), however, mentions that borings can vary from oval, circular, cross-shaped, and multi-sided polygons to extremely irregular ones. However, octopods are the only known predators to bore a distinctive ovoid

and bevelled hole (*Oichnus ovalis*). *Octopus* activity has been thoroughly neglected by palaeontologists because, having no skeleton, they are virtually unrepresented by body fossils. Thus, *O. ovalis* may be used to gain some insight into ancient octopod-feeding behaviour (Bromley, 1993; Harper, 2005; Pasini & Garassino, 2012; Klompmaker *et al.*, 2013, 2014; Klompmaker & Landman, 2021).

Crustaceans (stomatopods and decapods) are active predators of molluscs. The common injuries they cause to shells are chipping and peeling (*Caedichnus*) (Frey, 1987; Stafford *et al.*, 2015). Crustaceans produce still other injuries to shells, such as the holes (*Belichnus*) originating in gastropods and bivalves from blows produced by the chelipods of stomatopods (Geary *et al.*, 1991; Baluk & Radwanski, 1996). Successful chipping or peeling attacks can leave distinctive damage on the shells, whereas unsuccessful attacks are often repaired, leaving characteristic repair scars on the shell (Fig. 4E–4J) (Vermeij *et al.*, 1980, 1982; Vermeij, 1982; Allmon *et al.*, 1990; Turra *et al.*, 2005; Cintra-Buenrostro, 2007). As predation represents a major cause of death within marine ecosystems, shell-crushing predation by decapods has been shown to be an important process in the evolution of shells and crab claw designs (Vermeij, 1977; Vermeij *et al.*, 1981; Bertness & Cunningham, 1981; LaBarbera & Merz, 1992; Molinaro *et al.*, 2014). However, in cases other than predatory boreholes, it is often difficult to differentiate between *ante-* and *post-mortem* mechanical damage and *ante-mortem* predatory damage. Thus, it has proven difficult to attribute marginal shell damage in bivalves to predation by crustacean decapods, particularly crabs (Alexander & Dietl, 2003). However, a high incidence of death by breakage does not necessarily mean that breakage is evolutionarily significant. For example, if all attempts by predators to break shells resulted in the prey's death, the shell would be useless as a protection structure. The persistence of the prey populations facing such predation would occur not by virtue of the protective role of the shell, but in spite of it. The only way in which traits conferring resistance to cracking could spread through a population is by the survival and subsequent reproduction of individuals that have withstood onslaughts by shell-breakers. In other words, the evolution of anti-predatory protection is contingent upon predator failure (Vermeij, 1994). Repair scars also occur after non-lethal shell damage due to causes other than predation, e.g., impact by stones, or self-inflicted damage during predation and burrowing.

Durophagous shell breakage is abundant in the Neogene record (Martinell, 1989; Geary *et al.*, 1991; Baluk & Radwanski, 1996). Papp *et al.* (1947) described and illustrated numerous examples of predation by the peeling activities of decapods mainly related to Tortonian gastropod prey, highlighting that certain naticid gastropod species are more prone to shell repair than others. Martinell *et al.* (1982) reached similar conclusions for the marine Pliocene of Catalonia—the

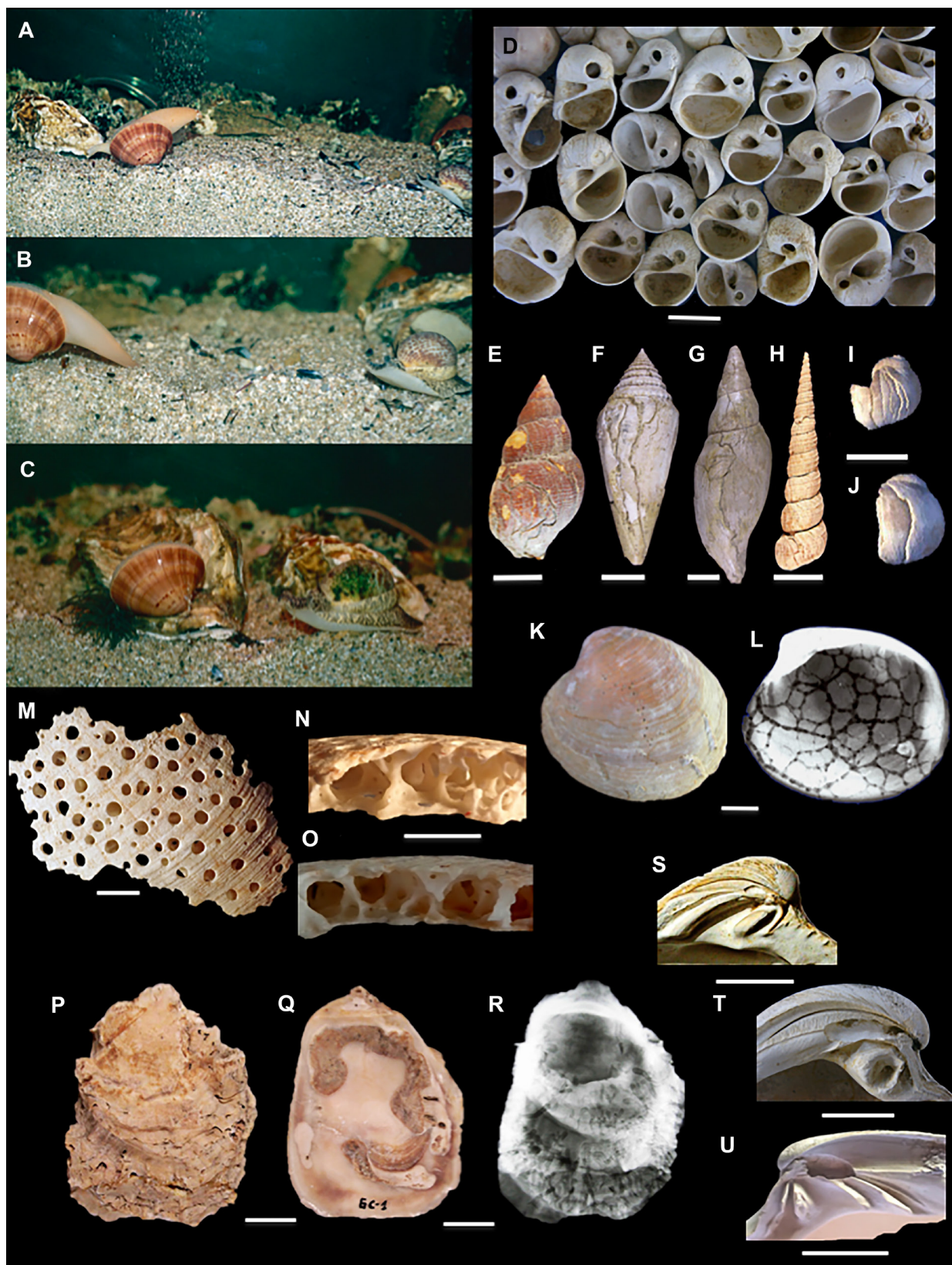


Figure 4. A–C, Attack sequence of the gastropod *Naticarius hebraeus* on the bivalve *Callista chione* in aquarium; D, *Cochlis pseudoepiglotina* shells showing naticid borings (*Oichnus paraboloides*), Pliocene, Roussillon, France; E–J, some predation scars made by crabs on Pliocene gastropods, and repaired; E, *Buccinum*; F, *Conus*; G, *Scaphella*; H, *Turritella*; I–J, *Cochlis*; K–L, *Arctica islandica* with *Entobia*; K, outer view of the shell; L, x-ray image to show the boring network in the shell; Pleistocene, Girona coast; M–O, *Entobia* boreholes in a shell fragment of a Recent *Glycymeris*, El Perelló, Tarragona; P–R, *Crassostrea rhizophorae* with *Maeandropolydora*; Q, clear mud blister; R, x-ray image to show the weakened interior of the shell; Pleistocene, Colombia; S–U, *Umbicnusus* in bivalve shells; S, *Venericardia planicosta*, Eocene, Île-de-France, France; T, *Pelecyora gigas*, Pliocene, Huelva, Andalucía; U, *Dosinia concentrica*, Pleistocene, Sarasota, Florida; scale bars = 1 cm.

gastropod species with the most shell repairs being those belonging to turritellids, naticids, and nassarids, while also highlighting that between 45% and 99% of the injured specimens exhibit more than one repaired scar.

Using the Plio-Pleistocene fossil record of crab predation on morphologically similar pairs of right-coiled and left-coiled gastropod species, it has been noted that the rare sinistral coiling snail promotes survival from attacks by right-handed crabs. Dextral busyconid and conid species typically have higher shell repair frequencies than sinistral ones (Dietl & Hendricks, 2006).

Parasitism. Bivalves and gastropods are known to be hosts for various parasites that live in their mantles or within their shells (Blake & Evans, 1973; Carroll & Finelli, 2015; Diez *et al.*, 2011; Sato-Okoshi *et al.*, 2012). Parasites, by definition, are harmful to their host to a certain degree. Negative effects can vary from minor metabolic changes to more important soft tissue damage and reduced shell strength. Parasites recognizable in the fossil record are those that leave traces of their activity on/in the shell of their hosts. Among them, the best known are sponges, polychaetes and bivalves (Cremonte, 2011).

Boring sponges belong to the Clionidae family and nowadays are the dominant endolithic sponges. *Cliona* is the Recent genus with the largest number of species, and specimens are capable of boring into a variety of substrates, including live and dead shells. They mainly parasitise epifaunal molluscan species (ostreids, pectinids, and mytilids) (Evans, 1969; Rützler & Rieger, 1973; Rützler, 2002; Rohde, 2005), and brachiopods (Bromley *et al.*, 2008). Sponges penetrate the periostracum forming 0.5–4 mm diameter circular holes on the surface and excavating a network of interconnected chambers or galleries with a series of globular chambers connected by canals and with external apertures, which appear as pores in shell surfaces (Fig. 4K–4O). In the fossil record they are registered under the ichnogenus name *Entobia* and its numerous ichnospecies. Apertures are invariably multiple, either linear or pepper pot-like in their distribution. The trace commonly appears parallel to the structure of the shells (De Groot, 1977; Bromley, 1970; Bromley & D'Alessandro, 1984, 1989; Garilli *et al.*, 2022).

Polychaete worms are active and widespread bioeroders, especially members of the family Spionidae. *Polydora websteri* Hartman is capable of boring into a variety of substrates, including coral, limestone, and live and dead shells (both bivalves and gastropods) (Voigt, 1975; Blake, 1971; Wargo & Ford, 1993). Its borehole is U-shaped, elongate and straight or sinuous, with the two parallel tubes separated by a central raised vane. The bore aperture has an eight-shaped outline. They commonly appear parallel to the layered structure of the shells. When the tubes reach

the inner part of the valve, the infested mollusc reacts by secreting conchiolin layers in their attempt to isolate the polychaete, which may contain mud and debris from polychaetes (mud blisters) (Fig. 4P–4R). The presence of mud blisters in fossil ostreid valves would be indicative that the drilling activity was carried out while the bivalve was alive, allowing identification of a parasitism relationship (Domènech *et al.*, 2008).

Various ichnotaxa have been described related to bioerosion caused by worms in general, *Maeandropolydora* and *Caulostrepsis* ichnogenera being associated with spionid polychaetes. They basically differ by the proximity (*Caulostrepsis*) or separation (*Maeandropolydora*) between the entrance and exit orifices and by the existence or not of a central raised vane along the route of the tunnels (Bromley & D'Alessandro, 1983). When infestation has occurred in a living shell, it is interpreted as parasitism and the producers are described as endoskeletozoans (Taylor & Wilson, 2002). As a reaction to the boring activity of the spionids, a thickening of the infected shell occurs. In the case of scallops, this thickening reduces their ability to swim, making them more accessible to predators (Mori *et al.*, 1985). Ishikawa and Kase (2007) proposed *Polydorichnus* as a new ichnogenus for spionid boreholes in gastropod shells. *Umbichnus inopinatus* Martinell *et al.* is another common ichnospecies located in an unusual but consistent position in the bivalve shells. The borehole is sack-shaped and is emplaced in and between the opposed hinge plates of the two shell valves, just beneath the umbo and dorsal to the hinge teeth. The size of the boring varies considerably, both within the same species of host, and according to the size of the host species. In large hosts, such as *Pelecypora islandicoides* Lamarck the borehole may reach a length of almost 2 cm (Martinell *et al.*, 1999) (Fig. 4S–4U). While the producer is still unknown, it has been attributed to some group of commensal polychaetes, probably spionids (Martinell *et al.*, 1999). The host bivalve is invariably a heterodont, 85 of the host species have been identified, distributed among the carditids, arcticids and venerids, from the Paleocene to the Recent (Martinell & Domènech, 2023). This borehole today appears to have a worldwide distribution, whereas in the geological record it has only been reported from the Mediterranean and northern Atlantic areas (Pereira *et al.*, 2009; Richiano *et al.*, 2012; Wisshak *et al.*, 2019; Martinell & Domènech, 2023).

Some solitary caryophylliid and flabellid scleractinian corals from the Pliocene of the Western Mediterranean exhibit epigenic groove-shaped bioerosional structures running along the surface of the thecae. Martinell and Domènech (2009) named them *Sulcichnus* and attributed them to the activity of polychaetes. Currently, similar structures are produced by the eunicid *Helmutneris flabellicola* Fage which maintains a commensalistic relationship with solitary corals. In the fossil record, *Sulcichnus* occurs associated to

shallow marine environments whereas their Recent counterparts are described for deep-marine corals. Martinell and Domènech (2009) interpreted this as a change in the environmental requirements of the coral/worm pair. *Sulcichnus* has been also found on Messinian rhodolith beds of the Western Mediterranean (Naimi *et al.*, 2021a, 2021b), though in this case the producer's commensalism habit must be discarded. Lithophagous bivalves can drill into the upper valve of large ostreids, mainly occupying the umbonal area. These borings (*Gastrochaenolites*) can become numerous, weakening the shell structure (Lauckner, 1983). Kleemann (2009) described *Gastrochaenolites* in corals as *Favia* from the Lower Eocene of Wascheberg (Austria) and *Montastrea* from the Miocene of Seggauberg (Austria), which have presented a slight S-like bending in the basal (old part) of the boring, indicating an adaptation of the boring axis to the change in direction of the host coral growth. It also could be evidence of parasitism of the lithophagous bivalve in the coral.

BIOEROSION VERSUS TAPHONOMY

Death

Death is the initial stage from which the processes of taphonomy begin. In other words, it is the initiation of the transition of an organic entity towards the fossil register (Emig, 2002).

As seen in the Predation section, some traces of bioerosion, e.g., complete drill holes produced by carnivorous gastropods, are obvious signs of the prey's death. Others, such as the traces of parasitism produced by spionid polychaetes (*Polydora*), can cause the death of the host even when the infestation only markedly weakens its shell. However, this fact is more complicated to verify since the activity of the spionids can continue to affect the host's shell after death. In addition, spionid polychaetes also pierce the shells of already dead molluscs, which would at this point simply be acting as a substrate.

The taphonomic processes on terrestrial vertebrate remains have received substantial attention both in the field and the laboratory (Weigelt, 1928, 1989; Behrensmeyer & Hill, 1980; Brain, 1981; Shipman, 1981; Andrews, 1990; Allison & Briggs, 1991; Lyman, 1994; Martin, 1999; De Renzi *et al.*, 2002; Fernández-Jalvo & Andrews, 2016). Mikuláš *et al.* (2006) erected three new ichnotaxa for bite and gnaw marks recognized in animal fossil hard tissues: *Machichnus* (shallow serial or subparallel grooves, interpreted as gnaw marks), *Nihilichnus* (circular or oval attack marks in compact bone), and *Brutalichnus* (strong impact marks, characterized by breaking the cortical bone towards its inner part). Jacobsen and Bromley (2009) proposed a new ichnogenus for tooth impressions, *Linichnus* (unique elongated toothed groove on bone surface). Taphonomic processes in marine aquatic

environments differ markedly from those in terrestrial environments. Due to the complexity of developing underwater observations, studies on the alteration of the skeletal remains of current marine vertebrates are very scarce and are basically concentrated on the study of the pre-burial stages of cetacean carcasses (Allison *et al.*, 1991; Smith & Baco, 2003; Rouse *et al.*, 2004; Haag, 2005; Dahlgren *et al.*, 2006; Amon *et al.*, 2013; Muñiz *et al.*, 2020). Unfortunately, fossil cetaceans with a significant part of the skeleton preserved are relatively rare and consequently the few documented cases of possible shark bite marks (*Nihilichnus*, *Linichnus* and/or *Brutalichnus*) prevalently correspond to fragmentary remains (Lambert & Gigase, 2007; Noriega *et al.*, 2007; Ehret *et al.*, 2009; Esperante *et al.*, 2009; Godfrey & Lowry, 2021), being challenging to differentiate whether they respond to predation or scavenging. Shark teeth are sometimes found associated along with cetacean remains. This could be direct evidence of shark attacks (Ehret *et al.*, 2009; Muñiz *et al.*, 2009). Bianucci *et al.* (2001) describe a specimen of *Astadelphis* (Cetacea, Delphinidae) from the Italian Pliocene with well-defined predation bite marks produced by a shark, including striae (*Linichnus*), sulci (*Nihilichnus*) and abrasions (*Brutalichnus*)—visible on the right dentary, on several ribs and on vertebrae—having lethal consequences.

Biostratinomy

Around 40 animal species can coexist within the carcasses of today's whales, including different types of fish (sharks, hagfish, etc.) and invertebrates (crabs, ophiuroids, molluscs, polychaetes, etc.), which feed on their remains (Dahlgren *et al.*, 2006). In the fossil record, traces of bone eater *Osedax* (Siboglinidae, Polychaeta) activity are found in whales, fish, and bird bones (Kiel *et al.* 2010, 2013; Higgs *et al.*, 2011). They have been described as *Clavichus ionassi* (Muñiz *et al.*, 2010, 2017; Höpner & Bertling, 2017). The presence of the scavenger cassid gastropod *Liracassis*, found in close association with Oligocene mysticete whale skeletons, indicate a novel trophic relationship (Nesbitt, 2005). Possible traces of osteophagous activity of potential decapod crustaceans have been reported in the scapula of a mysticete whale from the lower Pliocene of Southern Spain (Esperante *et al.*, 2009; Belaústegui *et al.*, 2017b). Clavate borings (*Gastrochaenolites* isp.) found in a cetacean skull from the Miocene of Tarragona provide the first evidence of bioerosion by pholadid bivalves on an autochthonous marine mammal skeleton (Belaústegui *et al.*, 2011, 2012a). Cione *et al.* (2010) reported *Gnatichnus* and *Radulichnus* grazing traces on penguin bones from the Argentinian Miocene. The echinoids and molluscs responsible for these two ichnogenes grazed on the substrate to feed upon exo- and endoskeletal algae and other microorganisms.

The remains of different organisms are transported in many ways. At the same time, similar organic

remains are transported in different ways in different environments. Most *post-mortem* transport of bottom-dwelling marine organisms depends on currents. In relatively shallow waters where currents are strong, a skeleton may be carried a considerable distance to its final burial site. Sorting or selection during the *post-mortem* transport is important because it influences the composition of the final fossil assemblage.

The left-right valve sorting by currents in bivalves varies depending on the size, the symmetry, and the perimeter of the shells (Martin-Kaye, 1951; Brenchley & Newall, 1970; Shimoyama & Fujisaka, 1993; Chattopadhyay *et al.*, 2013; Angseesing, 2020). Lever (1958) and Lever and Thijssen (1968) carried out extensive research on the transport behaviour of pierced valves (with *Oichnus*), concluding that there is a drill hole effect. The unpierced valves had a different behaviour compared to the bored ones, with differences depending on the number, location and diameter of the hole(s). The proportion of drilled/non-drilled valves, rather than strictly being an ecological factor, may have a taphonomic bias.

Although bioerosion is recognized to cause significant destruction of hard substrates, as of yet, few studies have assessed the loss of biological entities due to a taphonomic bias in relation to shells with predatory perforations (*Oichnus*). Some studies have assumed that drilled and non-drilled bivalve shells do not have the same preservation potential, indicating that valves with *Oichnus* are significantly weaker than those without them, considering factors such as transport, compaction, etc. (Roy *et al.*, 1994). On the contrary, Kowalewski and Hoffmeister (2004) tested whether drilling frequencies were independent of the taphonomic condition with shells from the Holocene beach ridges of Baja California and from mollusc assemblages of the European Miocene. These authors concluded that drilling did not affect the pre-burial survival of shells. The results from Kelley (2008) were similar, finding that the bias unfavourable to the specimens containing *Oichnus* was not significant in fossil shell beds examined—neither from the Miocene of Maryland nor the Plio-Pleistocene of Florida.

The transport of more delicate skeletons (e.g., scaphopods, polychaete tubs), which were also attacked by boring predators, can easily provoke their breakage across the borehole. In a quantitative analysis of predation in the serpulid polychaete *Ditrupa arietina* Müller from the Pliocene of Southeastern Spain, Martinell *et al.* (2012) observed that between 29% and 38% of drilled tubes had broken across the boring. These figures reveal, on the one hand, that information on the intensity of predation can be lost due to breakage while, on the other hand, that the skeletons could be more broken than would correspond to the same environmental conditions but without being affected by bioerosion.

On present-day beaches it has been observed that shells of certain species of bivalves appear with the

umbo worn away, giving an *Oichnus*-like structure. These are shells with prominent umbos (*Acanthocardia*, *Glycymeris*, etc.) that, once the animal is dead and the skeleton is subjected to sea currents, are eroded by friction at the umbonal end (Fig. 5A–5D). This breakage can be confused morphologically with naticid borings. However, its position at the umbonal end, and often its size, indicate with greater certainty that the hole is of mechanical origin (Martinell *et al.*, 2021a).

The transport of biological entities can also be carried out by living organisms. In marine ecosystems, hermit crabs play a major role in this process. Throughout the “life” of a gastropod shell, numerous hermit crabs may inhabit it, subjecting it to different ecological and environmental conditions than those of the gastropod’s own habitat. This process effectively extends the “shelf life” of the shell far beyond that of the initial gastropod (Walker, 1989). Shells inhabited by hermit crabs are susceptible to damage from predators that attack both gastropods and hermit crabs. These shells can also be affected by environmental agents, such as wave action and potential colonisation by boring organisms like sponges and spionid worms. Unlike living gastropods, hermit crabs cannot defend themselves against them (Bertness, 1982; Walker, 1989, 1992). Certain boring organisms—including foraminifera, sponges, polychaetes, barnacles or bryozoans—can penetrate shells of both living gastropods and those inhabited by hermit crabs, potentially producing visible holes and weakening the shell’s structure, and thus increasing its vulnerability to environmental damage. While skeletal dissolution and boring represent types of non-predatory damage, which is distinct from predatory damage, they can still weaken the shell’s integrity, making it more susceptible to environmental damage (Smyth, 1990). Effective shell availability to hermit crabs depends on the causes of gastropod death. The hermit crab population may utilize almost all shells, even those drilled and otherwise damaged, so that all habitable shells are exposed to the same types of taphonomic damage (Zuschin *et al.*, 2003; Ishikawa *et al.*, 2004; Stafford & Lighton, 2011). Hermit crabs can select which species’ empty shells they use. In the rocky supralittoral zone of the coast of Cabo Huertas (Alicante), practically out of the water, we observed monospecific groups of hundreds (almost 400) of *Cerithium lividulum* Risso shells occupied by the hermit crab *Clibanarius erythropus* Lastreille (Fig. 5E). Obviously, this habitat does not correspond to that of the living *C. lividulum*, a species who inhabits shallow but water-covered rocky bottoms. The result is the alteration of the molluscan taphocenosis and, thus, the loss or masking of palaeoecological and palaeoenvironmental information.

A variety of marine bivalves (clavagellids, mytilids, tridacnids, pholadids, myids, petricolids, arcids, and gastrochaenids) actively bore into invertebrate skeletal remains, using chemical and/or mechanical techniques, while producing club-shaped dwellings

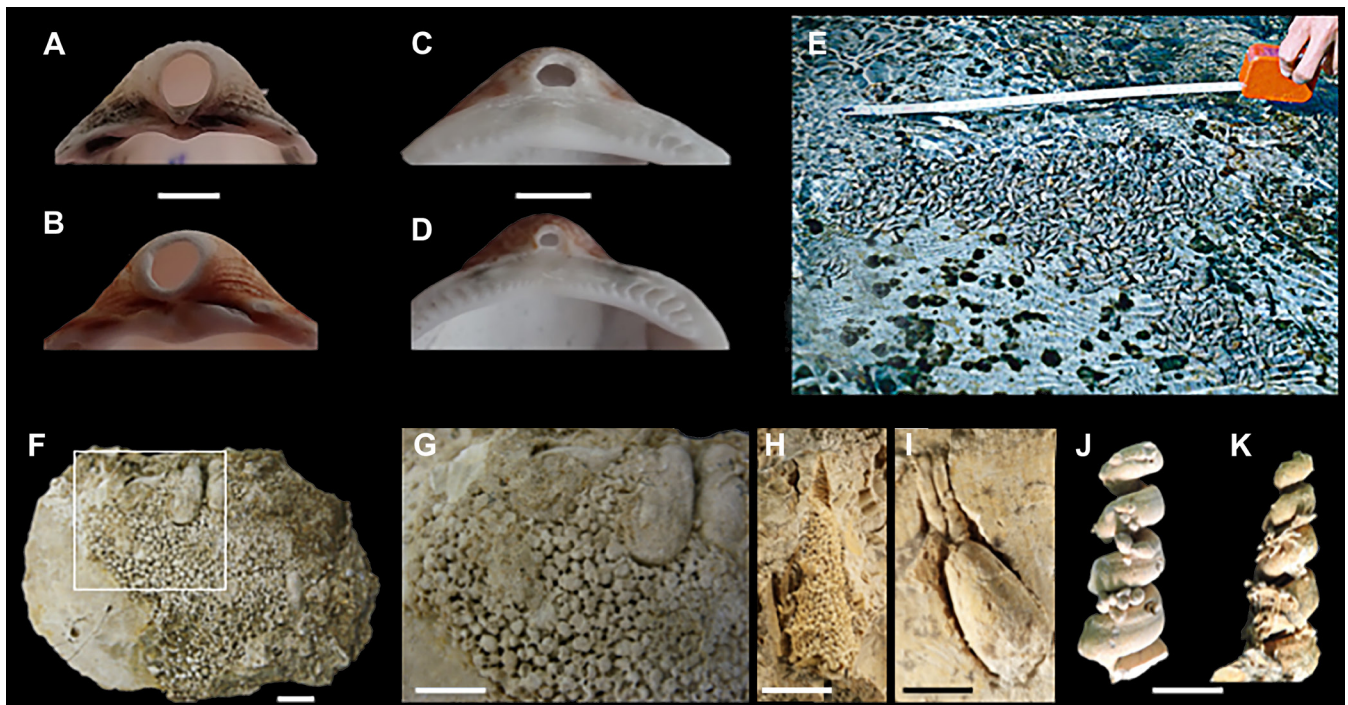


Figure 5. A–D, Bivalve shells with umbones “bored” by inorganic marine abrasion; A–B, *Acanthocardia*; C–D, *Glycymeris*; El Remolar beach, Baix Llobregat; E, *Cerithium lividulum* shells concentration by the hermit crab *Clibanarius erythropus*, Cabo de las Huertas, Alacant; F–G, inner mould of *Entobia* (detail in G) and some small moulds of *Gastrochaenolites* in a bivalve mould; H, *Entobia* mould completely covering a turrillid gastropod inner mould; I, mould of *Gastrochaenolites* including the siphons conducts; J, turrillid inner mould with small *Gastrochaenolites* moulds; K, turrillid inner mould with *Caulostrepsis* moulds. Specimens in F to K from Middle Miocene, Garraf, Barcelona; scale bars = 1 cm.

(*Gastrochaenolites*). Large and thick epifaunal bivalve shells (rudists, clams, oysters) have been described from the Mesozoic to the Recent, showing weakened shells due to perforations by lithophagous bivalves (Kleemann, 1994; Bromley, 2004; Villegas-Martín *et al.*, 2022). Several Miocene species of *Ostrea* from Patagonia (Argentina) have been reported with abundant *Gastrochaenolites*, usually in the upper valve (Farinati & Zavala, 2002; Mauna *et al.*, 2005; Domènech *et al.*, 2014; Romero *et al.*, 2017). Mauna *et al.* (2005) suggested that the absence of *Gastrochaenolites* in the inner side of the valves would indicate that the drilling activity of lithophages occurred while the oysters were still alive and, at the same time, that the oyster shells were quickly buried after death. Lopes (2011) reached similar conclusions through studying the bioerosive activity in *Ostrea puelchana* D’Orbigny and *Crassostrea virginica* Gmelin from the Pleistocene-Holocene of Rio Grande do Sul (Brazil).

The abundance of clypeasteroids in the fossil record is attributed to a variety of distinctive characteristics that enhance their preservation, in addition to the formation of mass aggregations. These characteristics include: (1) their tendency to gather in groups; (2) the strength of their skeletal structure; (3) their high mobility; and (4) their preference for inhabiting shoreface environments, which favour its redistribution and accumulation (Belaústegui *et al.*, 2012b). The dome-shaped tests of *Clypeaster* thus represented the commonest, largest,

and most stable hard substrates on an otherwise soft, shifting, sandy sea bottom, and have therefore been considered as benthic islands or preferential spots for settlement and growth of boring skeletozoans in the Neogene. *Gastrochaenolites* is one of the commonest bores in clypeasteroids fossil tests (Santos *et al.*, 2003b; Belaústegui *et al.*, 2013; Rahman *et al.*, 2015). The distribution of modern wood-boring bivalves is mainly controlled by salinity, temperature, and availability of wood (Lavigne, 1999). Teredinids thrive in exclusively marine conditions but can survive in brackish waters with salinities over 10 ppt and grow at higher rates in warmer water conditions.

Wood-boring teredinid and pholadid bivalves produce clavate or club-shaped borings, which are collectively assigned to the ichnogenera *Teredolites* (Kelly & Bromley, 1984) and *Apectoichnus* (Donovan, 2018), being present in the stratigraphic record since the Jurassic. *Teredolites* is taphonomically informative because it provides evidence for the timing and position of wood prior to burial. Heise *et al.* (2011) observed that (on average and after two years on the sea bottom) the 74% loss of original volume in modern wood samples was due to teredinid colonization. *Teredolites* have been found both in *in situ* composite wood substrates (Bromley *et al.*, 1984; Savrda *et al.*, 1993), in individual allochthonous fossil log-grounds, and as ghost log-grounds (*Teredolites* liberated from their original wood-substrates) (Pickerill *et al.*, 2003;

Kříž & Mikuláš, 2006). Because of the relative ease with which logs are transported in aquatic systems, *Teredolites* found in allochthonous trunks is by far the most common situation (Radwanski, 2009; Donovan *et al.*, 2009; Monaco *et al.*, 2011; Villegas-Martín & Rojas-Consuegra, 2011; Kumar *et al.*, 2011; Villegas-Martín *et al.*, 2012; Leszczynski, 2018; Serrano-Brañas *et al.*, 2019; Belaústegui *et al.*, 2020a).

Fossildiagenesis

Resedimentation of preserved elements is a biostratinomic process, whereas the reelaboration of those elements (named reworking in many papers) is a fossildiagenetic process (Fernández López, 1984a, 1984b).

The reelaboration of fossils is commonly associated with “poor” preservation, as processes such as breakage, abrasion, corrosion, encrustation and boring could affect them after exhumation and transport. Reelaborated fossils commonly correspond to robust skeletal remains (shells, bones, wood) or to fossil moulds that were eroded out from a sedimentary rock. Reelaborated fossils may remain buried in either younger or older strata, depending on their taphonomic journey. The character of reworked fossils is easily deduced through their bioerosive structures when these affect skeletal moulds (Akpan *et al.*, 1982; Fernández López, 1984b). The analysis must be more precise when the perforations directly affect the shell, as for example in the case of corals, oysters, rudists or clypeastoroids, where it must be discerned whether the traces are contemporary with the skeletal remains or were produced during the reworking process (Donovan *et al.*, 2010; Villegas-Martín *et al.*, 2022). On modern beaches, limited by sedimentary sequences, it is not unusual to find a mixture of present-day fauna and recently drilled fossils (Donovan, 2012; Donovan *et al.*, 2014; Lopes, 2024).

A paradigmatic case would be that of underwater deposits with perforated mollusc shells from the marine Pleistocene of the NE of the Iberian Peninsula (Domènech & Martinell, 1982; Martinell & Domènech, 1981, 2023; Molinu, 2015; Belaústegui *et al.*, 2020b). The typical fauna of cold waters appears exposed in the canyons of the marine platform, where they could have been used as a substrate by today’s macro- and micro-borers. Distinguishing between the activity that took place in the vital or biostratinomic phases from that which would have occurred in the last stages of the fossildiagenetic phase is relevant for the interpretation of the taphonomic assemblage, given the different environmental conditions that would correspond to different drillers/encrusters. It is therefore hard to deduce the drilling sequence and, consequently, to know the actuation order for each producer.

Ekdale *et al.* (1984) defined lithoturbation as a process of successive episodes of bioerosion-sedimentation-cementation-bioerosion. As such, lithoturbation

is a perfect example of the relationships among biological, biostratinomic and diagenetic processes. The deciphering of the successive events permits the interpretation of a depositional environment as well as the reconstruction of the fossilisation path. Belaústegui *et al.* (2018) describe and interpret a multi-episodic marine shell accumulation in the Pi Gros Miocene outcrop (Garraf, NE Iberian Peninsula) where both bioerosional, biostratinomic and diagenetic processes act successively and repeatedly on skeletons, moulds and substrate (Fig. 5F–5K). Most notably, bioerosion traces due to different endoskeletozoans gave clues regarding the sequence. Boreholes produced in shells by sponges, polychaetes and bivalves were simultaneous, but the presence of semi-endoskeletal dwellings and isolated casts implies episodes with higher sedimentation rates. Furthermore, clusters of intersecting clavate-boring *G. dijugus* proved the occurrence of different bioerosive episodes. Due to diagenesis, fossil molluscs were preserved as internal and external moulds, while traces were preserved as positive casts. This site provided an excellent example of how both destructive (dissolution) and constructive (cementation of boring casts) diagenetic processes may result in the exceptional preservation of bioerosion trace fossils, which at the same time explains a complete multi-episodic sequence. Taxonomic, ichnologic and taphonomic approaches have allowed the interpretation of the depositional environment and the fossilisation succession.

Romero Baena *et al.* (2022) also present an extraordinary case of *Gastrochaenolites* preservation in opaline silica. The borings were produced by pholadid bivalves in volcanic tuff pebbles deposited in a marine subtidal environment after a tsunami event (Oligocene, Ica, Peru). The mineralization affected both the pholadid shells and the borings and allows for the precise interpretation of the diagenetic environment.

SOME REFLECTIONS, BY WAY OF CONCLUSIONS

Bioerosion studies are in good health. A handful of researchers from all over the world have been specializing in this field and both basic ichnotaxonomic results and those applied to palaeontology and geology have been increasing in recent years. This means that the future of this field can be faced with optimism.

Bioerosion is a destructive-constructive process of palaeobiological information present in all the Phanerozoic marine ecosystems. Bioeroders modify substrates (rocky, skeletal, or wood) through destruction, but at the same time build new (palaeo) biological information that needs to be interpreted carefully.

Macroperforations may not be as spectacular as other groups of ichnofossils (e.g., dinosaur footprints) or most body fossils, but their study provides valuable information about the biodiversity and ecology of

past marine ecosystems, and moreover about the taphonomic processes which affected the organic remains.

The ichnotaxonomic composition of rocky-shore assemblages changed through the Phanerozoic (*Tyanites-Gastrochaenolites-Entobia*). Macrobioerosion in rocky shores is mainly used to characterize unconformity surfaces, which is of great interest from a geological point of view. Nevertheless, I believe that more detailed studies on the possible relationships among the rock borers themselves are needed as they could contribute new information on ancient rocky ecosystems. A few articles have been published in this sense and there is still much terrain to explore.

The study of depredation through the investigation of bioerosion traces in skeletal remains is seminal in order to see tendencies, coevolution, etc. in the geological past, but it is also important to take into account the fact that we are dealing with organisms that do not always follow the same behaviour. For example, when we are registering naticid traces on the shell of their victims, it is worth considering that they do not always need to bore to obtain food. In this case, predation activity is not reflected in the palaeontological record, although it had existed.

The drill hole effect regarding the transport of the bivalves by marine currents is a taphonomic bias that must be taken into account in palaeoecological studies, whether they are approached from a global or analytical alternative.

From the taphonomical point of view, is relevant to assess the response of bioeroded skeletal remains—especially those affected by endolithic drillers—being weakened shells, which are more fragile when transported by sea currents and which, consequently, can fragment more easily even with weak currents.

The study of bioerosional processes in present-day coastal environments and beaches is very useful for interpreting biostratinomic processes (transport, fragmentation, etc.) in ancient, shallow environments. Beaches are authentic changing laboratories of taphonomic processes that marine dynamics modify depending on the characteristics of the currents, their intensity and direction, and which in turn modify the components of biological entities. We can see the results in the thanatocoenoses accumulated on today's beaches (from the shoreface to the backshore), usually riches in mollusc remains. They are particularly evident after storms or episodes of rough seas, but they also occur under calmer conditions.

Donovan *et al.* (2010) have reported that “Reworked macrofossils are rare and fascinating components of the record that can preserve evidence of unique *postmortem* histories.” I agree with the concept of fascinating, but I do not think that reworked (*sensu* reelaborated) macrofossils are rare but, rather, that they have not yet received enough attention. Until this point, we have only taken into account all the present-day planetary rocky shores built by sedimentary rocks

that are potentially rich in fossils. Marine erosion is occurring to a greater or lesser extent, and it produces reelaboration of fossils that fall into the present sea and can be used as settlement sites for many organisms, encrusters, and bioeroders; we find them in the present-day beaches. Over more than 600 Ma of regular erosion in rocky shores, the production of reelaborated entities must have been really important. I think this is a fascinating field in which the study of bioerosion can make a significant contribution to our knowledge of the past and present.

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