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The Ordovician System in Estonia

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Ordovician rocks are widespread in the Baltoscandian region. The main distribution area of Ordovician strata in the East European Platform extends from the Baltic Sea islands in the west to the vicinity of Moscow in the east and from the Gulf of Finland in the north to Belarus and Poland in the south. In the northern part of this area, in the eastern coastal region of the Baltic Sea, beds are exposed in the magnificent sections of the Baltic–Ladoga Klint, in other coastal and river bank sections, old and new limestone quarries and open-cast pits of northern Estonia and northwestern Russia. Good accessibility of geological exposures and excellent preservation of fossils and sedimentary structures attracted the attention of investigators already in the early 19th century when the strata were described and figured by O. M. L. v. Engelhardt (1820), W. Strangways (1821), E. Eichwald (1825) and others. The researchers have been interested in the characteristic Cambrian to Middle Ordovician succession, represented by several distinctive rock units (the Cambrian Blue Clay, phosphatic brachiopod coquina, *Dictyonema* argillite, dark green glauconite sandstone and a wide variety of limestone units above).

The main features of the Ordovician stratigraphy were first outlined by F. Schmidt in his thorough monographic paper of 1858. The general pattern of his geological map, presented in the same volume, is well recognized in the modern geological maps of the bedrock of Estonia. The generally simple geological structure of the area, with almost horizontal strata, only 2–5 m/km dipping to the south, results in nearly latitudinal orientation of the outcrop belts of the Ordovician stages in northern Estonia (see the geological map of Estonia on the back cover of the present volume).

The main part of the Ordovician succession in northern Estonia is composed of limestones, with some intercalations of kukersite oil shale concentrated mainly in the Kukruse Stage. Only the basal Ordovician strata comprise a relatively thin succession of clastic sediments – sandstones, argillites and clays of the Pakerort and Varangu stages, overlain by the glauconitic sand- and siltstones of the Hunneberg and Billingen stages. The transition from the terrigenous to carbonate rocks in the Billingen Stage is marked by the appearance of calcareous interbeds in the siltstones, which grade into the first limestone/dolomite unit, the Toila Formation. The appearance of the first representatives of the numerous characteristic Middle Ordovician fossil groups is recorded in the same transition interval or in the overlying Volkhov Stage.

The Ordovician limestone succession in Estonia and adjacent areas begins with cold-water carbonates deposited in a sediment-starving shallow marine basin. The sedimentation rates have increased upwards. Changes in sedimentation rates in the calcareous main part of the Ordovician succession are in obvious correlation with the carbonate production rates. The corals make their first appearance in the Upper Ordovician, and the first carbonate buildups can be recorded in about the same interval, emphasizing a striking change in the overall character of the palaeobasin.

Generally the change in the type of sedimentation and the character of biofacies is ascribed to a gradual climatic change resulting from the northward drift of the Baltica Palaeocontinent from the temperate climatic zone to the (sub)tropical realm (Nestor & Einasto 1997). During the Middle and Upper Ordovician the climatic change resulted in an increase in the carbonate production and sedimentation rate on the carbonate shelf, whereas the deposition pattern was controlled by the accommodation space available there.

The details, but also the problems of the Ordovician geology in the subsurface area in central and southern Estonia were first revealed only in the 1950s. A large number of drill cores, obtained in the course of an extensive drilling programme in the 1950s–1980s, revealed a marked difference between the stratigraphic successions in the outcrop area and southern Estonia. As a result of the comparison of the eastern Baltic and Scandinavian successions, the concepts of the structural-facies zones (by Männil

1966) or confacies belts (by Jaanusson 1976) were introduced for the Ordovician of Baltoscandia. As the term "confacies" is unique (being exclusively used for the Ordovician of Baltoscandia), а different terminology that is widely applied in newer publications has been introduced by Harris et al. (2004; Fig. 1). The micropalaeontological and macrofaunal studies of the core sections have also revealed the distinctive biogeographic differentiation pattern, characteristic of the Ordovician rocks (e.g. Männil 1966; Männil et al. 1968; Männil & Meidla 1994; Meidla 1996). Although the biofacies pattern is described for the eastern Baltic area, the facies zonation of the entire Baltoscandian area is still imperfectly known. The

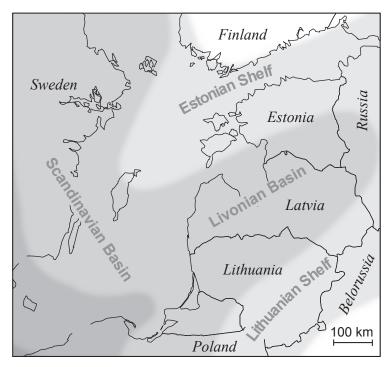


Fig. 1. Post-Tremadocian Ordovician facies zonation.

seismic investigations of the Baltic Sea area performed in the last decades (Tuuling 1998; Tuuling & Floden 2013 and references therein), but also new detailed (micro-)palaeontological investigations (*e.g.* Tinn & Meidla 2001) might produce valuable new information in this field.

The total thickness of the Ordovician in Estonia varies from 70 to 180 m, being greatest in central and eastern Estonia and considerably less in the outcrop area, as well as in the southwestern mainland of Estonia.

Several correlation problems still persist in the Ordovician of Estonia, due to marked biofacies differences between northern and southern Estonia. In part, they are discussed also in a recent monographic overview of Estonian geology (see Heinsalu & Viira 1997, Meidla 1997, Hints 1997, Hints & Meidla 1997 in Raukas & Teedumäe 1997). During the last years the application of the stable carbon isotopic zonation (Ainsaar *et al.* 2010) has opened new opportunities to solve still persisting problems in regional stratigraphy.

The term "Stage", first employed by Bekker (1921), has become the principal category in the chronostratigraphic classification of the Ordovician System in Estonia. The development of the stratigraphic classification of the Ordovician strata in Estonia, from the "beds" (Schichten) by Schmidt (1858) to the stages in modern meaning is documented in detail in Männil (1966), Rõõmusoks (1983) and Rõõmusoks et al. (1997). The term "Ordovician" was introduced for Estonia by Bassler (1911). A number of regional series and subseries for the Ordovician System in Estonia and neighbouring Russia were brought into use by Schmidt (1881) and several subsequent authors. Raymond (1916) applied the traditional American three-fold subdivision of the Ordovician System to this particular area, but this classification was subjected to repeated changes until 1987. Also the terms "Oeland Series", "Viru Series" and "Harju Series" have been widely used as a basic classification for the Ordovician System of the area since the 1950s (introduced by Kaljo et al. 1958 and Jaanusson 1960 in a nearly recent meaning). The subseries have been introduced as well (see Männil & Meidla 1994 and Nõlvak et al. 2006 for a summary), but they are very rarely used today and the well-established framework of the Ordovician System has in fact replaced the regional suprastadial units in publications. The modern threefold classification of the Ordovician System (IUGS 2013) was first used for the Estonian succession by Webby (1998) and is presented here in detail (Fig. 2).

System							Correlation of the formations (<i>members</i>)			
SV 0	Series	Stage	Graptolite zones	Conodont zones and subzones	Stage & index, substage		NW Estonia	northern Estonia	central Estonia	southern Estonia, western Latvia, NW Lithuania
Silur.	Lland.	Rhud.	Coronograpt. cyphus. ?	Distomodus kentuckyensis	JUURU G ₁₋₂		Hilliste Ta	amsalu Varbola	Õhne	Õhne
		rnantian	Metabolo. persculptus Metabolograptus		PORKUNI	[Saldus	Saldus Kuldiga
	-	Hi	extraordinarius Dicellograptus anceps	ordovicicus	PIRGU	U	Årina Adila	Årina Adila	Adila Halliku	Jelgava
	pper		Dicellograptus complanatus		F ₁ c	L	Moe	Moe	Moe Jo	onstorp
			Pleurograptus linearis	?	VORMSI F	U U	Kõrgessaare Saunja	Kõrgessaare Saunja	Tudulinna Saunja	Fjäcka Saunja
		Katian		Amorphognathus superbus	F ₁ a RAKVERI E	E	Paekna Rägavere	Paekna Rägavere	Mõntu Rägavere	Mõntu
		k	Dicranograptus clingani ?		OANDU D _{III}	-		Hirmuse	Hirmuse Varik	u Westerner Westerner Westerner
				KEILA D ₁₁		Vasalemma			≥ Plunge Blidene	
		Diplograptus foliaceus Nemagraptus gracilis			HALJALA	U	Kahula	Kahula	Kahula	Adze
R			Amorphognathus tvaerensis	C _{III} -D _I KUKRUSE	L	Tatruse	Tatruse	Tatruse Dreimani	Dreimani	
Ordovician		_	? Gymnograptus	Pygodus anserinus ⁹ /8	C _{II}	U	Pihla ↓↓↓↓↓↓↓↓↓↓ Kõrgekallas	Viivikonna Kõrgekallas	Kõrgekallas	Taurupe
0		-	linnarssoni Pseudoamplexograpt. distichus	Pygodus serra $\begin{bmatrix} 6_5^7 \\ 4_1 \\ 3 \end{bmatrix}$	C ₁ c LASNAMÄ C ₁ b	L GI	Väo	Väo	Väo (upper) Stirnas	Stirnas
	dle	rriwilian	Pterograptus elegans	Eoplacognathus suecicus	ASERI C ₁ a			Kandle	Rokiškis	Segerstad
Middle	Mid	Da	Nicholsonog. fascicul. Holmograptus lentus	Eoplacognathus pseudoplanus	KUNDA B _{III}	U M	Pakri	Loobu	Loobu (lower)	Baldone <u>Šakyna</u>
		Daping.	Didymograptus hirundo	Yangtzeplac. crassus Lenodus variabilis Balt. norrlandicus Paroistodus originalis Baltoniodus navis Baltonio. triangularis	VOLKHOV B _{II}	L U M L		Sillaorg (lower)	Toila	Kriukai
Lower		- <u></u>	1 en ugi . uppi өлітиниз	Oepikodus evae Prioniodus elegans	BILLINGE B _i b HUNNEBER	N RG		Leetse	Leetse	Zebre
	ower	Iremadocian	Hunnegrap. copiosus, Adelograp. murrayi, Adelograptus hunnebergensis	Paroistodus proteus Paltodus deltifer C angulat C rotund	B ₁ a VARANGU A ₁₁₁	J	Varangu	Varangu		
		Trema	Rhabdinopora flabell. anglicum-R.f. multith. Rhabdinopora flabel-	C. angulatC. rotund. Cordylodus lindstroemi	PAKER-	U	Türisalu	Türisalu		
Cambrian	Furongian		liformis norvegicum Rhabdinopora flabelliformis sociale	Cordylodus intermed. Cordylodus proavus Cordylodus andresi	ORT A _{II}	L	Kallavere	Kallavere	Kallavere	Kallavere

Figure 2. Ordovician stratigraphy of Estonia. Graptolite zonation according to Kaljo & Vingissaar, 1969, Kaljo *et al.*, 1986, Männil, 1976, Resheniya..., 1987, Männil & Meidla, 1994, Nõlvak *et al.*, 2006, conodont zones according to Kaljo *et al.*, 1986, Meidla, 1997 and Männik in Nõlvak *et al.*, 2006. Numbers in the column of the conodont zonation correspond to the conodont subzones as follows: subzones of the *Baltoniodus norrlandicus* Zone: 1 – *Trapezognathus quadrangulum* Subzone, 2 – *Lenodus antivariabilis* Subzone; subzones of the *Pygodus serra* Zone: 3 – *Eoplacognathus foliaceous* Subzone, 4 – *Eoplacognathus reclinatus* Subzone, 5 - *Eoplacognathus robustus* Subzone, 6 – *Eoplacognathus protoamosus* Subzone, 7 – *Eoplacognathus lindstroemi* Subzone; subzones of the *Pygodus anserinus* Zone: 8 – *Sagittodontia kielcensis* Subzone, 9 – *Amorphognathus inaequalis* Subzone; subzones of the *Amorphognathus tvaerensis* Zone: 10 – *Baltoniodus variabilis* Subzone, 11 – *Baltoniodus gerdae* Subzone, 12 – *Baltoniodus alobatus* Subzone.

In relation to the definition of the GSSP for the base of the Ordovician System in the Green Point section, Newfoundland (Remane 2003), a revision of the traditional position of the Cambrian–Ordovician boundary at the base of the Pakerort Stage in Estonia turned out to be necessary. According to conodont data, the systemic boundary in the northern Estonian sections lies a few metres higher than previously suggested, i.e. in the middle of the Pakerort Stage, within the Kallavere Formation (Puura & Viira 1999). The lower boundary of the Silurian System has traditionally been drawn at the base of the Juuru Stage. The most recent results on the stable carbon isotopic chemostratigraphy disagree with this viewpoint and emphasize that the evidence behind this traditional solution is not convincing, arguing for a higher position of this boundary in the regional succession, within the Juuru Stage.

Main features of the chronostratigraphic (stage) classification of the Ordovician System were outlined already by Männil (1966). Only minor changes in stage nomenclature have been made in the later decades: the Ceratopyge Stage has been renamed the Varangu Stage (Männil 1990), the Latorp Stage replaced by the Hunneberg and Billingen stages (Hints *et al.* 1993) and a new unit, the Haljala Stage, has merged the former Idavere and Jõhvi Stages (following Jaanusson 1995 and Nõlvak 1997) that were difficult to distinguish outside northwestern Estonia. Hints & Nõlvak (1999) brought the concept of boundary stratotypes ("golden spike") into the Estonian stratigraphy, proposing a stratotype – the Pääsküla outcrop – for the lower boundary of the Keila Stage that also marks a faunal change in the succession. However, as stratigraphic hiatuses on the stage boundaries are very common in northern Estonia (and all remarkable faunal changes are usually related to hiatuses), wide usage of this concept for the stage boundaries in this area still seems rather complicated.

Graptolites are rare in the carbonate succession of Estonia. The Scandinavian graptolite zones are usually adopted for the correlation charts, although their correlation to the local succession is mainly based on indirect evidence as the local graptolite record is well resolved in some intervals only (the intervals of the Pakerort–Varangu and Uhaku–Kukruse stages (Kaljo & Kivimägi 1970, 1976; Kaljo *et al.* 1986; Männil 1966, 1976, 1987; Nõlvak *et al.* 2006; Fig. 2). Conodont zonation, however, has been elaborated in detail and is of very high resolution for the Lower and Middle Ordovician (see Fig. 2).

The elaboration of lithostratigraphic classification of the Ordovician rocks was initiated by Orviku (1940), who proposed the lithostratigraphic subdivision for the upper Middle Ordovician. This approach was widely accepted by subsequent authors and led to description of a very substantial number of formations and members and compilation of a series of detailed correlation charts approved by the Interdepartmental Stratigraphic Committee of the former USSR (Resheniya... 1965, 1978, 1987 and a related paper by Männil & Rõõmusoks 1984). The last version of such a formal correlation chart (the edition of 1987) was, in a slightly emended form, published also in English, in the series of the IUGS publications (Männil & Meidla 1994). The correlation chart in Fig. 2 contains some recent improvements compared to this publication, the most recent ones being introduced by Ainsaar & Meidla (2001) and Nõlvak *et al.* (2006). Some more modifications of the Ordovician correlation charts for Estonia have been published by Hints *et al.* (1993) and Nõlvak (1997). The composition and textures of the Ordovician carbonate rocks and main differences between the confacies belts were summarized by Põlma (1982 and references therein).

The monographic studies on the Ordovician palaeontology started already in the 19th century. After the comprehensive review on the Ordovician and Silurian strata (in modern meaning) by Schmidt (1858 and several subsequent monographic papers), a number of important monographic papers were published by F. B. Rosen, W. Dybowski, A. Pahlen, G. Holm, A. Mickwitz, O. Jaeckel, J. H. Bonnema and R. F. Bassler. The tradition of palaeontological investigations on the Ordovician material of Estonia was continued by A. Öpik (1930, 1934 and others) and, later on, by the recent generation of palaeontologists. Monographs and extensive monographic papers have been published on the Ordovician brachiopods, corals, stromatoporoids, chitinozoans, scolecodonts, ostracods, conodonts, etc. Summaries on the palaeontological investigations on virtually all fossil groups recorded from the Ordovician of Estonia are published in the monograph *Geology and mineral resources of Estonia* (Raukas & Teedumäe 1997; http://sarv.gi.ee/geology/).

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